

Effect of intermittent gravity wave activity on the dynamics in the mesosphere during winter

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Introduction

- Gravity waves (GWs) are the major contributor to vertical coupling of atmospheric layers by distributing energy and momentum throughout the whole atmosphere.
- GWs have horizontal wavelengths of tens to hundreds of kilometers so that GWs are mostly parameterized in global circulation models.
- GW distributions are mainly based on GW source parameterizations, or on specific functions or observed GW fields [Šácha et al., 2016; Lilienthal et al., 2017].
- Effects of different spatial and temporal GW distributions have to be studied to reproduce a more realistic circulation in the middle atmosphere.

MUAM – Middle and Upper Atmosphere Model

- Primitive equation 3D grid point model [Pogoreltsev et al., 2007]
- horizontal resolution: $5^\circ \times 5.625^\circ$
- Upper boundary: 160 km (log-p); $\Delta z = 2.842$ km
- Nudging of ERA-Interim zonal mean temperature below 30 km
- GW parameterization: linear scheme with multiple breaking levels
- GW initialized at 10 km (1 cm s⁻¹ vertical velocity perturbation)



Fig. 1: Horizontal resolution of MUAM.

Model setup: GW distribution

1. Spatial distribution:

- Instead of using the GW distribution based on the observed E_{pot} data we calculated the momentum flux (MF) distribution by applying midfrequency approximation according to Ern et al. [2004]:

$$F_h = 3\rho \frac{f}{N} E_{pot}$$

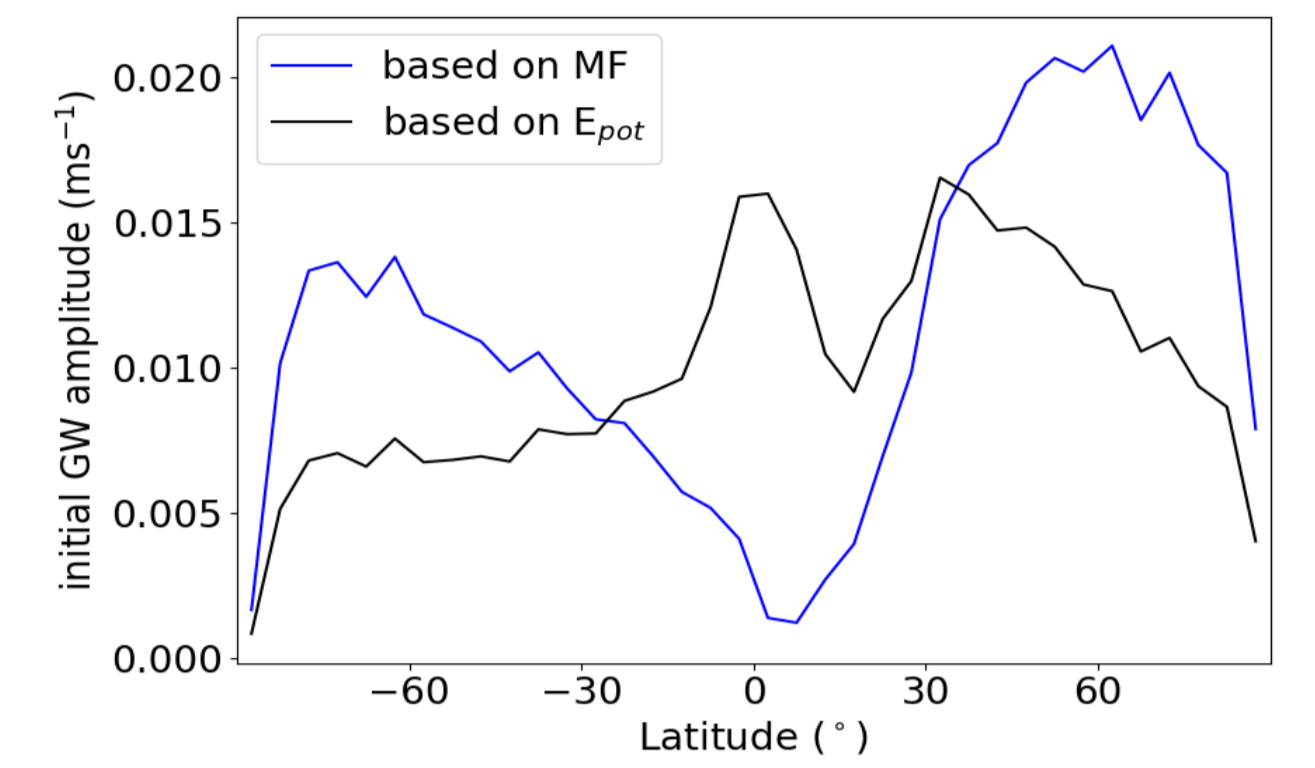


Fig. 2: GW vertical wind amplitudes for the simulation based on the GPS RO E_{pot} (dotted line) and MF (solid line) data for January conditions.

2. Temporal distribution:

- Real GW MF is unevenly distributed [Hertzog et al., 2012]: many GWs with small and only a few ones with large MFs.
- Create temporal variation of GWs by multiplying GW amplitudes with randomly generated numbers with mean values of 1 cm s⁻¹, 0.55 cm s⁻¹ and 0.275 cm s⁻¹.
- Distributed linearly and polynomially

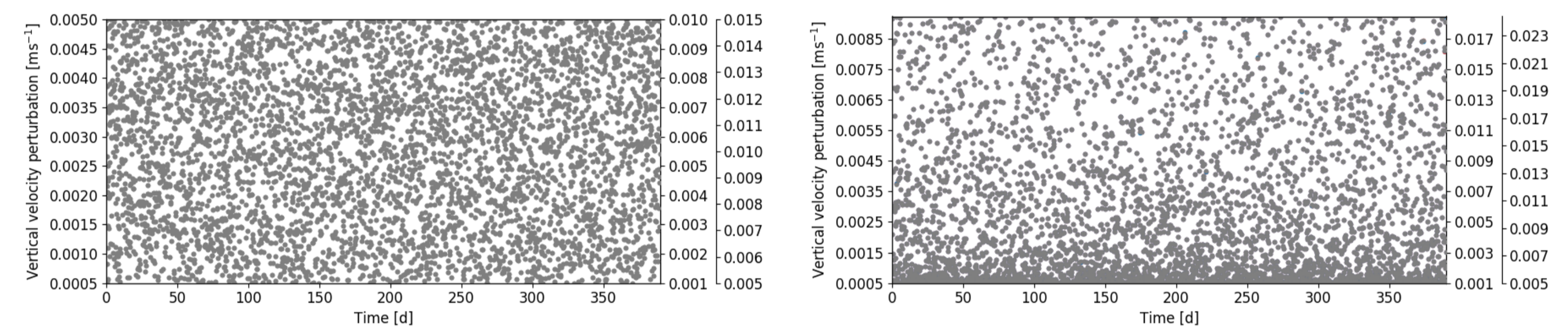


Fig. 3: Linear (left) and x^3 (right) distribution of the random numbers representing the variation of the vertical velocity perturbation with mean values of 1 cm s⁻¹, 0.5 cm s⁻¹ and 0.275 cm s⁻¹.

Results: Spatial GW distribution

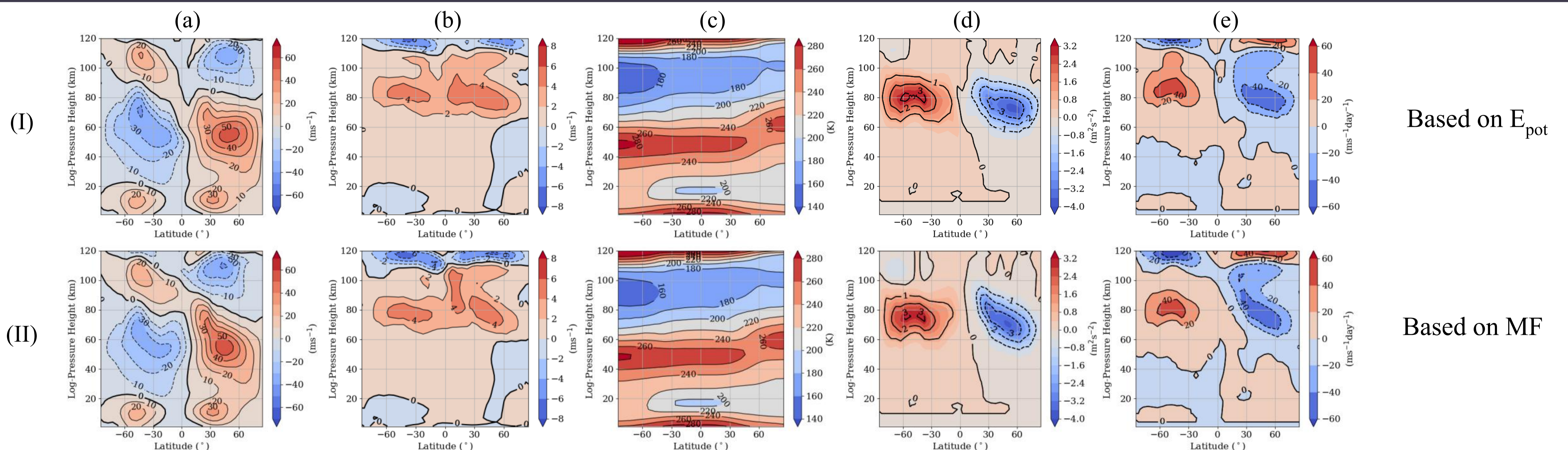


Fig. 4: January zonal and monthly mean of the (a) zonal wind [ms⁻¹], (b) meridional wind [ms⁻¹], (c) temperature [K], (d) GW fluxes [m²s⁻²] and (e) acceleration through breaking GWs [ms⁻¹day⁻¹]. GW parameterization is based on E_{pot} (I) and MF (II) data.

→ Tilt of northern hemispheric mesospheric jet towards lower latitudes with increasing height and of southern hemispheric jet towards the northern hemisphere.

Results: Temporal distribution of GW amplitudes

- Increasing height of the wind reversal with decreasing mean value of vertical velocity perturbation (not depending on temporal GW distribution).
- Decreasing strength of the mesospheric jet with increasing mean value of vertical velocity perturbation (smallest for polynomial distribution).

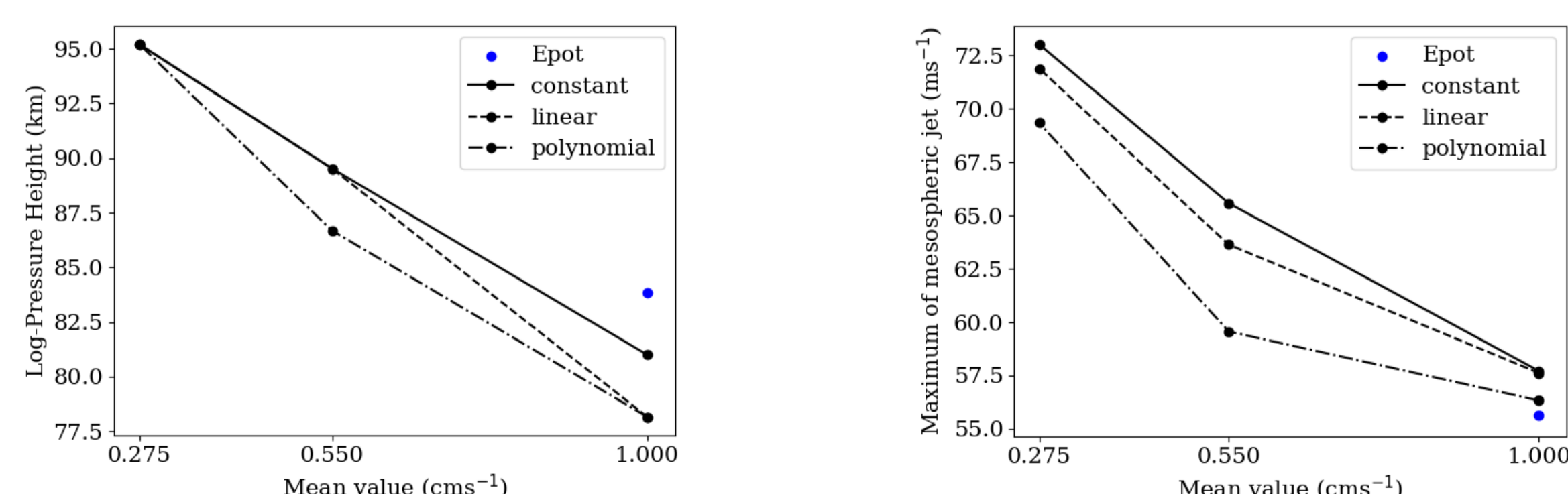


Fig. 5: Overview of the height of the wind reversal at 52.5°N (left) and the strength of the mesospheric jet (right) for each simulation.

Conclusion and outlook

- GW MF distribution leads to a more realistic background field (compare Figs. 4(Ia), 4(IIa) and Fig. 6).
- Further optimization through:
 - a composite of the simulations based on a mean of 1 cm s⁻¹ with a minimum of 0.05 cm s⁻¹.
 - adjustment of the phase speed.
 - implementation of a non-linear GW parameterization

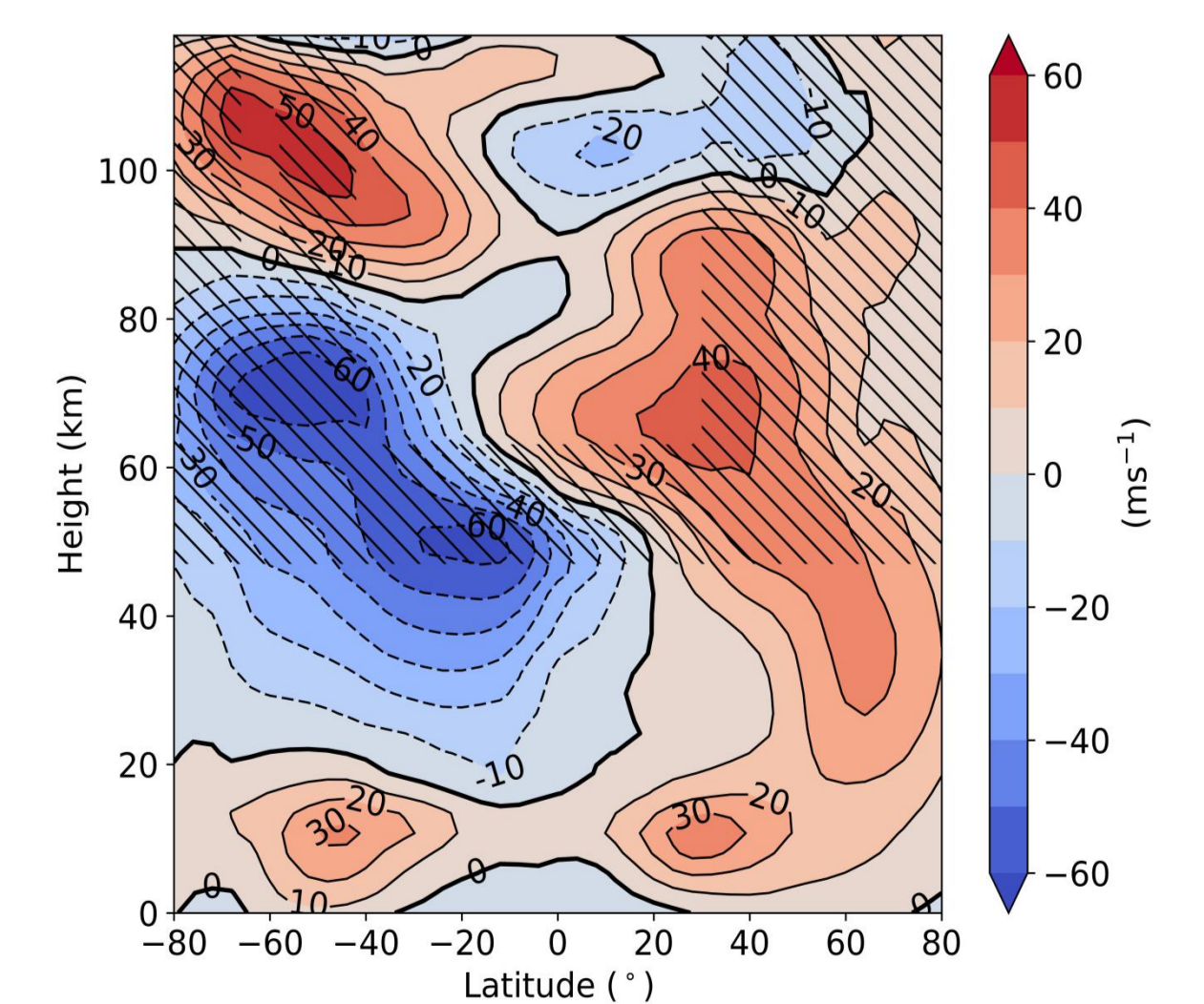


Fig. 6: URAP latitude-height plot of the zonal wind

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