



# Airborne remote sensing of cloud droplet number concentration using synergetic passive solar and microwave radiation measurements

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## 1 Motivation

- Radiative forcing of clouds depends on cloud properties ( $\tau$ ,  $r_{\text{eff}}$ , N, LWP, LWC)
- Same LWC, different N  $\rightarrow$  different albedo (Twomey-Effect)<sup>[2]</sup>
- Cloud inhomogeneity (LWC, N) of marine shallow cumulus mostly not covered by satellite observations

## 2 Campaign NARVAL-II

### Next generation Aircraft Remote-sensing for Validation studies

- Using the High Altitude and Long Range Research Aircraft (HALO) of DLR
- Subtropical and ITCZ region
- 08 August – 30 August 2016
- 10 Flights, 94 hours, 74000 km
- Investigate maritime shallow cumulus

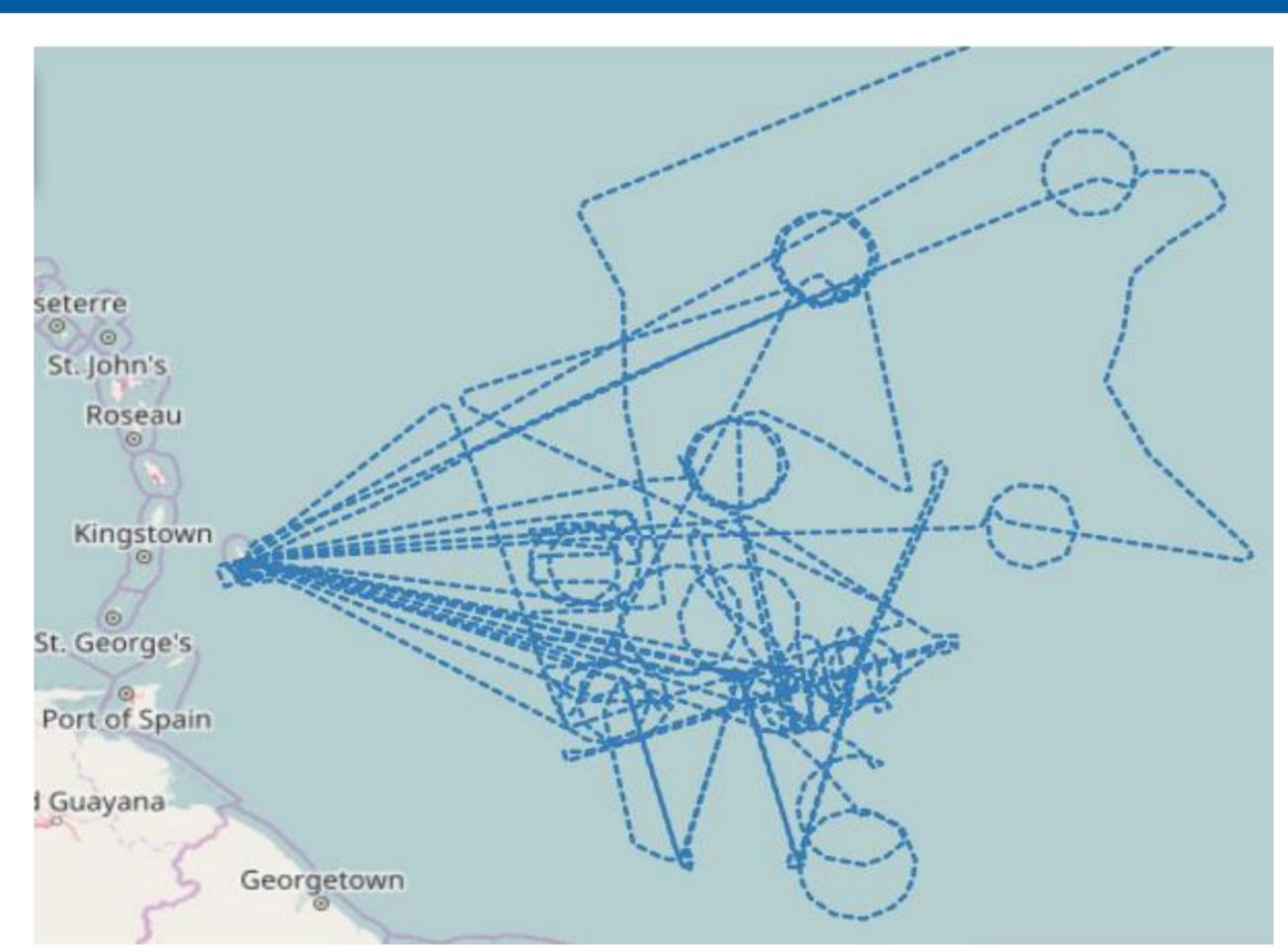


Fig. 1: Illustration of research flights with HALO during NARVAL-II.

## 3 Instrumentation

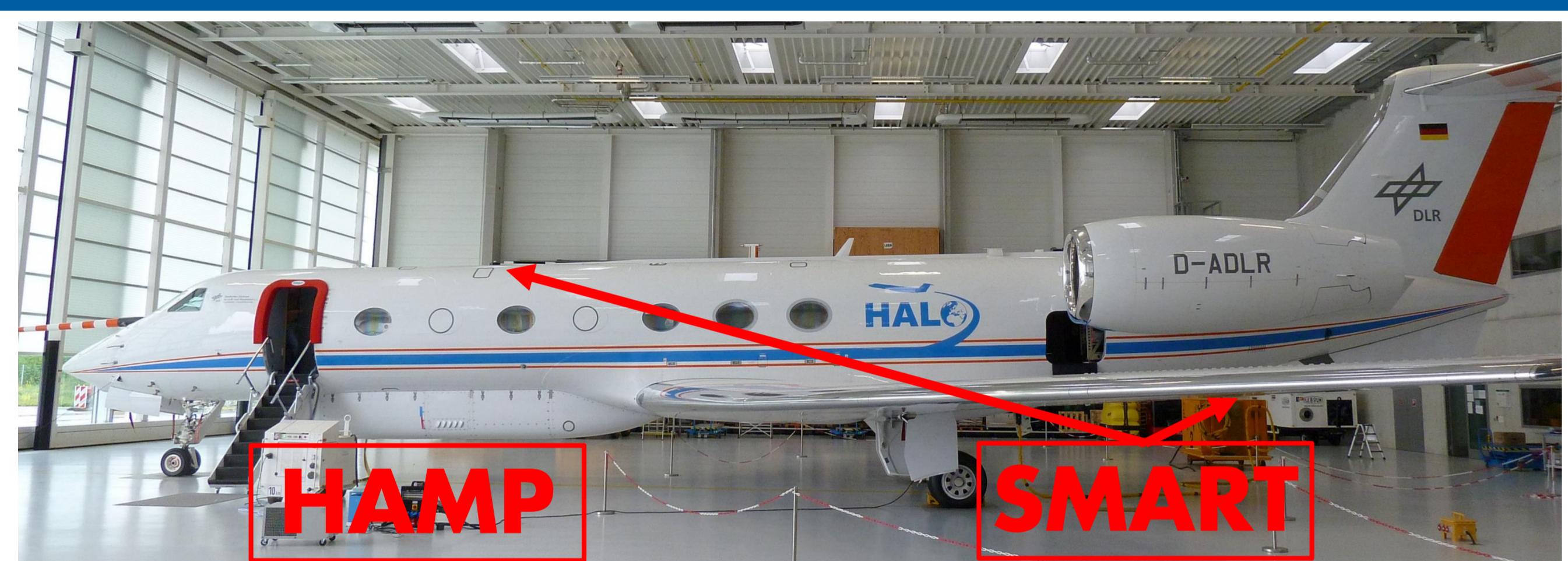


Fig. 2: Location of the HAMP radiometer + radar and SMART optical inlets.

### SMART-Albedometer

Spectral Modular Airborne Radiation measurement system - Albedometer



- Spectral solar irradiance ( $F_{\downarrow}^{\uparrow}$ ) and radiance ( $I^{\uparrow}$ ) with 2° FOV
- Spectral range: 300–2200 nm
- Temporal resolution 2 Hz
- Retrieved quantities: optical thickness, effective radius, therm. phase, cloud radiative forcing

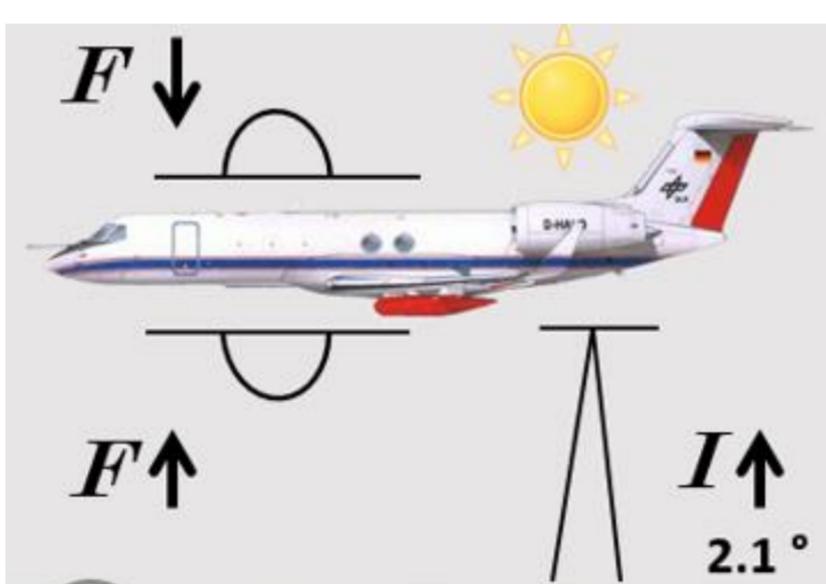


Fig. 3: Upper and lower optical inlets of SMART-HALO and measurement geometry.

### HAMP

HALO Microwave Package

- 23-channel microwave radiometer and cloud radar
- Retrieval of temperature and humidity profiles, LWP and hydrometeor properties

## 4 Method

### Retrieval of $\tau$ and $r_{\text{eff}}$

- Use of spectral cloud top reflectivity
- Cloud mask algorithm
- Applying iterative retrieval with radiative transfer model

### Retrieval of Liquid Water Path (LWP)

- From HAMP measurement

### Calculating Droplet Number Concentration<sup>[1],[3],[4]</sup>

- Assuming: adiabatic cloud profile (non-precipitating)
- Spherical cloud droplets
- Only liquid water clouds
- Geometrical thickness  $z$  from LIDAR
- Avoid Condensation rate assumptions

$$N_d = \frac{\sqrt{(10)}}{4 \cdot \pi \cdot k} \cdot \sqrt{\frac{2 \cdot LWP}{\rho_w \cdot r_{\text{eff}}^5} \cdot \tau}$$

## 5 Results from Flight 19.08.2016

### Retrieval of Cloud Droplet Number Concentration

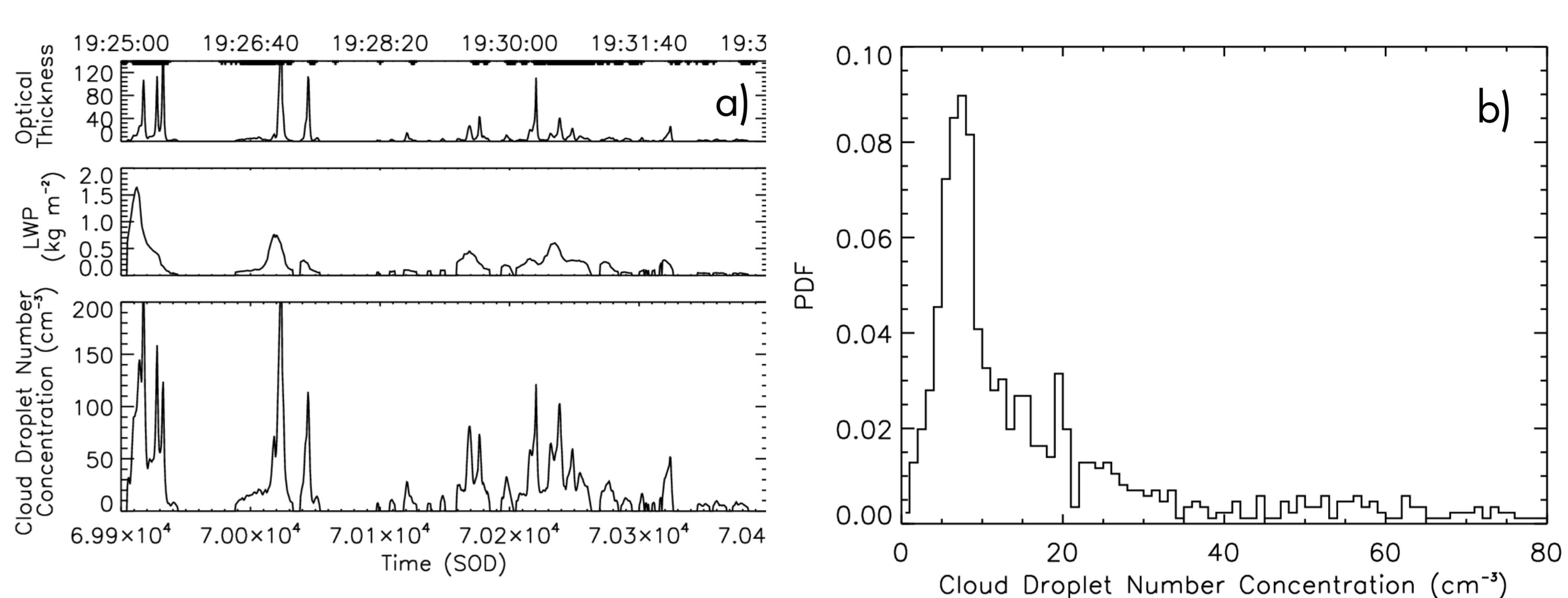


Fig. 4 a: Time series of optical thickness, liquid water path and retrieved cloud droplet number concentration. Fig. 4b: Probability density function of cloud droplet number concentration.

### Influence of Cloud Microphysics

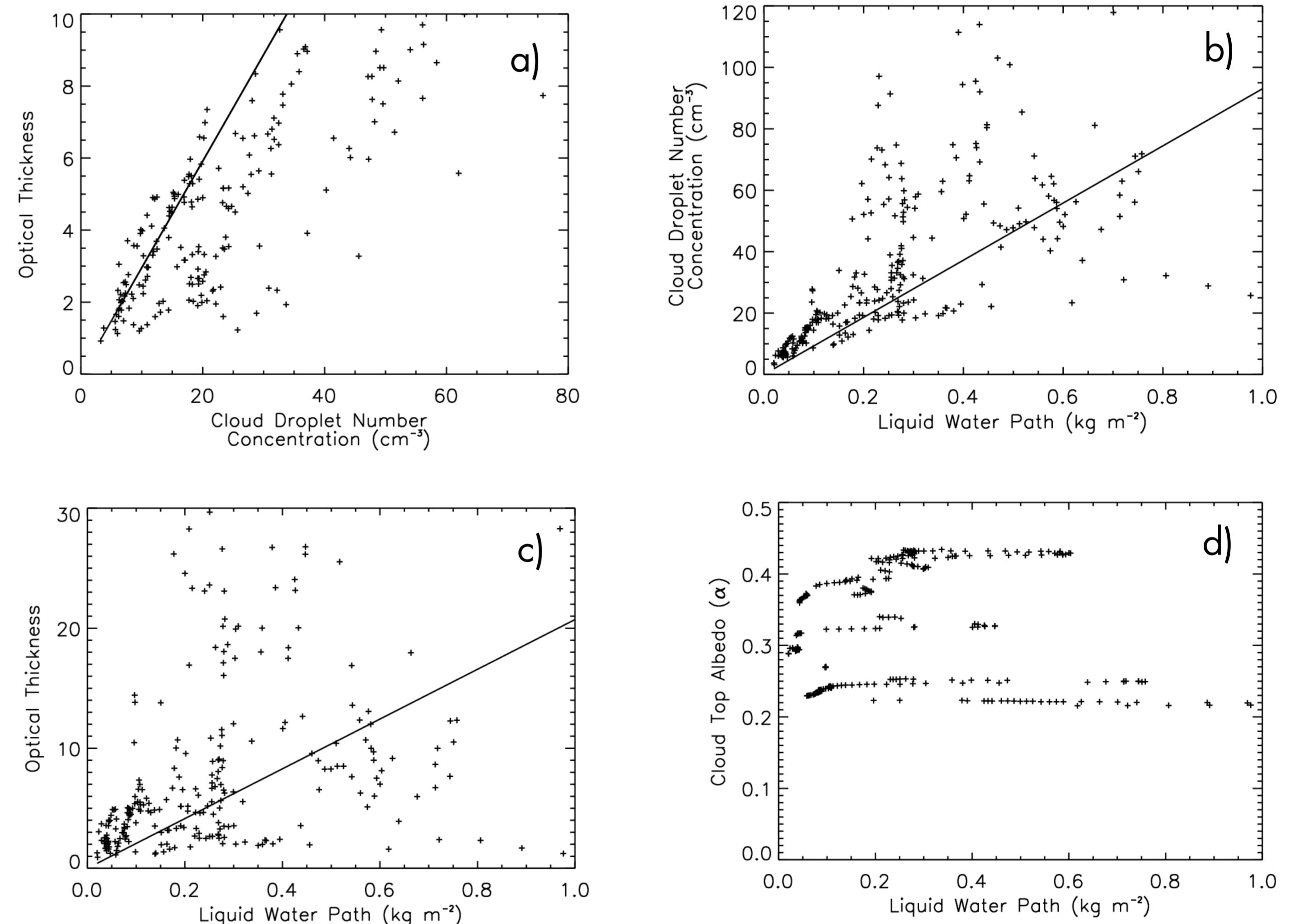


Fig. 5 a: Influence of cloud droplet number concentration on optical thickness. Fig. 5 b,c, d: Influence of liquid water path on cloud droplet number concentration, optical thickness and cloud top albedo.

### Influence of Cloud Distribution

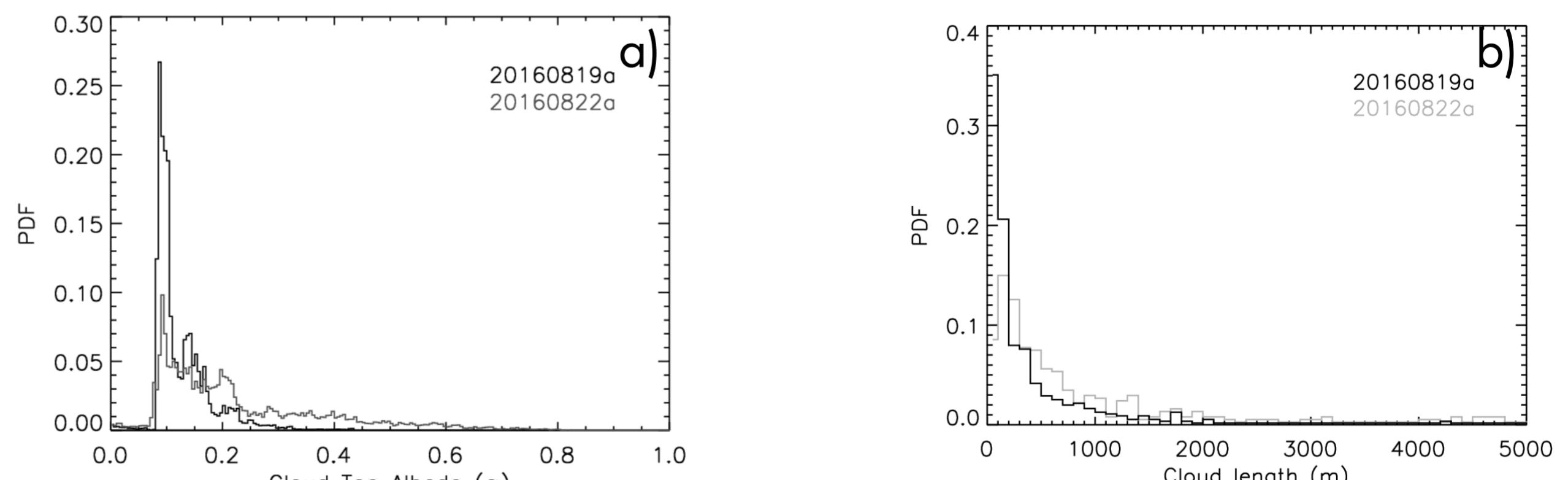


Fig. 6 a: Probability density function of cloud top albedo as a measure of total cloud cover for 19. and 22. of August. Fig. 6 b: Probability density function of cloud length for the same days.

## 6 Conclusion

- Synergetic retrieval of cloud droplet number concentration is possible
- Cloud top albedo depends on cloud inhomogeneity (cloud fraction and size)
- Planned satellite comparison / validation  $\rightarrow$  investigate (sub-pixel heterogeneity)

### References

- [1] Brenguier, J.-L. et al., 2000: Radiative properties of boundary layer clouds: Droplet effective radius versus number concentration, *J. Atmos. Sci.*, 7, 803-821, doi: 10.1175/1520-0469(2000)057<0803:RPOBL>2.0.CO;2
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- [3] Bennartz, R., 2007: Global assessment of marine boundary layer cloud droplet number concentration from satellite, *J. Geophys. Res. Atmos.*, 112, doi: 10.1029/2006JD007547
- [4] Zeng, S. et al., 2013: Study of global cloud droplet number concentration with A-Train satellites, *Atmos. Chem. Phys.*, 14, 7125-7134, doi: 10.5194/acp-14-7125-2014